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Planktonic Foraminifera biostratigraphy at the Jerissa area, (CES section in north-western Tunisia), and the impact of the Cenomanian-Turonian Oceanic Anoxic Event (OAE-2) on their assemblages

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During the late Cenomanian—early Turonian Oceanic Anoxic Event (OAE-2), organic-rich black shales were deposited worldwide. This event coincides with the strongest eustatic Uppermost Cenomanian transgressive period. Induced by this sea-level rise, the Oxygen Minimum Zone (OMZ) impinged onto the South Tethyan margin where the organic-rich Bahloul Formation indicates the record of OAE-2 in central Tunisia as well as north eastern Algeria.

At the CES section in the Jerissa area, a closer sampling along a 24m thick interval in this formation allows a high resolution biostratigraphic analysis based on planktonic Foraminifera. It shows that this event was extended throughout the upper part of the *Rotalipora cushmani* Zone and across the *Whiteinella archaeocretacea* Zone. There, five subzones are recognized which are correlated with those of the Pueblo global boundary stratotype section and point (GSSP). In the upper part of the *R. cushmani* Zone, the *Dicarinella algeriana* subzone coincides with the onset of the organic-rich deposits. The *W. archaeocretacea* Zone is divided into three subzones: the *Globigerinelloides bentonensis*, *D. hagni and Heterohelix moremani* subzones coincide with maximum TOC contents (up to 1.91%). At the top of the Bahloul Formation where the marls and limestones become impoverished in organic matter the *Helvetoglobotrunana helvetica* Zone starts.

Immediately below the Bahloul Formation in the upper part of the Fahdene Formation, samples have yielded diversified foraminiferal assemblages, containing more than 70% planktonic taxa. They are dominated by unkeeled surface dwellers such as of *Gl. bentonensis* associated with rare keeled deeper water dwellers (e.g. *Rotalipora cushmani*, *R. greenhornensis*, rare *R. montsalvensis*) and weakly-keeled lower photic zone dwellers (e.g. *Praeglobo-truncana stephani P. gibba*, *P. delrioensis*) and primitive dicarinellids. This assemblage points to an outer shelf to bathyal and

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oxygenated marine environment. Such an assemblage is preserved in organic-poor (0.2% TOC) light shales and marls with 30% CaCO₃ content. The first occurrence (FO) of the unkeeled surface dwellers of W. archaeocretacea is observed at this level. There, also the benthic foraminifera are diversified and dominated by bathyal species such as of Cassidella, Gavelinella, Lenticulina, Frondicularia, Gaudryina, Buliminella and Textularia.

From the upper part of the R. cushmani Zone (D. algeriana subzone), where the Bahloul Formation starts, the black shales are dominant and enriched in organic matter. There, an increase in the Planktonic/Benthic ratio is observed. It coincides with a major decline in the number of keeled Rotalipora species resulting from the disappearance of R. montsalvensis followed by the extinction of R. cushmani and R. greenhornensis (the LO of R. cushmani at -93.90 + 0.2 Ma). This particular coincidence of events is correlated in both Tethyan and Boreal realms. Throughout the middle part of the Bahloul Formation, where the Whiteinella archaeocretacea zone is developed, a drastic change in the benthic assemblage is also observed. This bioevent coincides with a positive shift of heterohelicids (Heterohelix moremani, H. reussi, H. aff. pulchra and H. navaroensis), then guembelitriids (Guembelitria cenomana, G. albertensis). Especially, heterohelicids are considered as low oxygen tolerant surface dwellers. Throughout the Gl. bentonensis subzone, their frequencies increase is related to the OMZ setting. Later, across the *D. hagni* subzone, guembelitriids mark a positive shift. They are considered as eutrophic surface dwellers. Their thriving coincides with maximum TOC values (2% TOC) and CaCO₃ contents (up to 80%). It may be related to the OMZ expansion.

The FO of the keeled $Helvetoglobotruncana\ helvetica\ (-93.29\pm0.2\ Ma)$ was noted at the top of the Bahloul Formation followed by $Marginotruncana\ spp.$ above through the Kef Formation where the TOC values again decrease.